



Provenance History of Detrital Diamond Deposits, West Coast of Namaqualand, South Africa

David Phillips¹, Jeff W. Harris² and Mike C.J. de Wit³

¹The University of Melbourne, Parkville, Australia, dphillip@unimelb.edu.au.

²University of Glasgow, Glasgow, Scotland, Jeff.Harris@glasgow.ac.uk.

³Tsodilo Resources Ltd, Toronto, Canada, mdewit@tsodiloresources.com.

Introduction

The West Coast of Namaqualand in South Africa hosts extensive detrital diamond deposits, but considerable debate exists as to the provenance of these diamonds. Some researchers have suggested that the diamonds were sourced from the erosion of >115 Ma orangeites located along the western part of the Kaapvaal craton and transported by the palaeo-‘Karoo’ river system, to be deposited in the Atlantic ocean near the current outlet of the Olifants river (e.g., de Wit, 1993). Other workers have argued that the majority of Namaqualand diamonds originated from recent erosion of Permo-Carboniferous Dwyka glacial deposits (ca.300 Ma), with their ultimate source being pre-Karoo kimberlites in the interior of the Kaapvaal craton (e.g., Moore and Moore, 2004).

In the current study, we analysed clinopyroxene inclusions extracted from a suite of detrital diamonds from coastal deposits along the Namaqualand coast in an effort to constrain the provenance of these diamonds.

Previous Work

Previous ⁴⁰Ar/³⁹Ar analyses of clinopyroxene inclusions extracted from diamonds sourced from the Orapa kimberlite produced a range of apparent ages from the time of kimberlite eruption to ~100 Ma older than this event (Phillips and Harris, 2008). As the ages were significantly younger than inferred diamond formation ages (>1.0 Ga), the ⁴⁰Ar/³⁹Ar results were interpreted to indicate significant argon diffusion to the diamond/inclusion interface zone prior to kimberlite emplacement, with >90% of radiogenic argon lost during inclusion extraction. Therefore, although this approach does not constrain the time of diamond formation, it provides a novel means of estimating (maximum) kimberlite eruption ages and is useful for constraining the provenance of detrital diamond deposits (Phillips and Harris 2009).

In a subsequent study, we analysed clinopyroxene inclusions from detrital diamonds collected from the west coast of Namibia and Namaqualand, using the ⁴⁰Ar/³⁹Ar laser probe dating method (Phillips and Harris, 2009). These analyses yielded maximum ages for the time of emplacement of source kimberlites/orangeites, mostly younger than ca.300 Ma. These data were used to infer diamond provenance from <300 Ma kimberlites/orangeites located on the Kaapvaal craton. However, these results were too imprecise to ascertain whether the diamonds were sourced from ca.85 Ma Group I kimberlites, >110 Ma orangeites, or possibly from both (or other) sources.

Results

In the current study, clinopyroxene inclusions were extracted from 35 Namaqualand detrital diamonds. ⁴⁰Ar/³⁹Ar ages were then determined using a new generation multi-collector ARGUSVI mass spectrometer system capable of ultra-high precision analyses (e.g. Phillips and Matchan, 2013; Phillips et al. 2017). Large inclusions were step-heated in two increments, whereas smaller inclusions (<200 µm) were fused in single analyses. Of the specimens analysed, 30 of 35 produced measurable signals, giving apparent ages ranging from 121.5 ± 3.3 Ma to 668.7 ± 4.9 Ma (2σ). Only six clinopyroxene inclusions yielded (maximum) apparent ages older than 300 Ma, with the majority (22 of 30) being <260 Ma. No clinopyroxene inclusions produced ages <100 Ma.

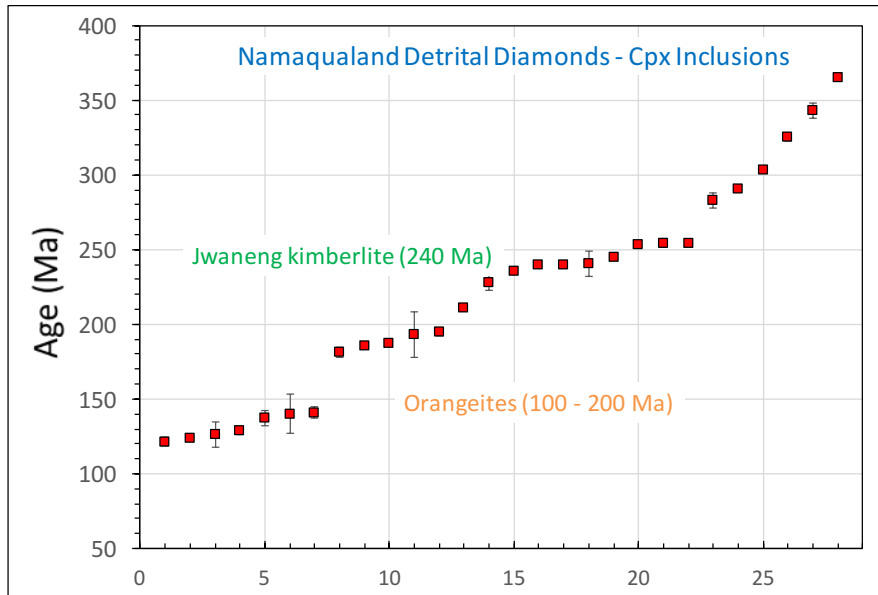


Figure 1. $^{40}\text{Ar}/^{39}\text{Ar}$ apparent ages determined on clinopyroxene inclusions extracted from detrital diamonds sources from Namaqualand coastal deposits (error bars are $\pm 2\sigma$). These ages represent maximum estimates for the time of source kimberlite/orangeite eruption due to the retention of up to ~10% of pre-eruption argon by the inclusions (see Phillips and Harris, 2009 for details).

Conclusions

The $^{40}\text{Ar}/^{39}\text{Ar}$ clinopyroxene results confirm that the vast majority of Namaqualand diamonds were sourced from post-Dwyka kimberlites/orangeites. These could include Late Cretaceous Group I kimberlites (80 – 90 Ma), Cretaceous/Jurassic orangeites (110 – 200 Ma) and/or early Triassic kimberlites (e.g. ~240 Ma Jwaneng kimberlite). However, it is noteworthy that none of the inclusions yielded ages typical of most Group I kimberlites (80 – 90 Ma). Although these ages are necessarily maxima, the data imply dominant diamond provenance from Cretaceous/Jurassic orangeites rather than Cretaceous Group I kimberlites. This conclusion accords with palaeo-drainage reconstructions in the area, which indicate a change in drainage in the mid-Cretaceous from a southwesterly directed palaeo-‘Karoo’ fluvial system to the current westerly directed Orange river drainage basin (de Wit, 1993).

References

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